

Eulerian Simulation of Dispersion in the ASL

Write a program to calculate the mean concentration field $\bar{c} = \bar{c}(x, z)$ downwind from a continuous crosswind line source¹ at $x = 0, z = h_s = 0.46$ m in the horizontally-uniform, neutrally-stratified atmospheric surface layer. Assume \bar{c} is the solution of the advection-diffusion equation

$$\bar{u}(z) \frac{\partial \bar{c}}{\partial x} = \frac{\partial}{\partial z} \left[K(z) \frac{\partial \bar{c}}{\partial z} \right] \quad (1)$$

with mean windspeed and eddy diffusivity profiles

$$\bar{u}(z) = \frac{u_*}{k_v} \ln \frac{z}{z_0}, \quad (2)$$

$$K(z) = \frac{k_v}{S_c} u_* z \quad (3)$$

($k_v = 0.4$ is the von Karman constant, u_* is the friction velocity, z_0 is the surface roughness length and S_c is the turbulent Schmidt number; the x -axis points along the mean wind direction, and z , as usual, is the vertical axis).

Let $J=1\dots N$ label a set of gridplanes along the vertical axis, and $I=1\dots$ a set of gridplanes spaced along x . For this assignment use the following implicit discretization, derived in class using the control volume method:

$$\begin{aligned} \Delta z \bar{u}_J [\bar{c}_{I,J} - \bar{c}_{I-1,J}] &= \Delta x K_{J+\frac{1}{2}} \frac{\bar{c}_{I,J+1} - \bar{c}_{I,J}}{\Delta z} \\ &- \Delta x K_{J-\frac{1}{2}} \frac{\bar{c}_{I,J} - \bar{c}_{I,J-1}}{\Delta z}. \end{aligned} \quad (4)$$

This is a “marching” problem in the sense that the x (alongwind) axis is a “1-way” axis — there is no mechanism for material to travel against the wind. Therefore given the profile $\bar{c}_{I-1,J}$ for all J at location $I-1$, we can compute the profile at I by rearranging to obtain a set of neighbour equations linking $\bar{c}_{I,J}$ at all J (this discretization results in an implicit algorithm). The notation $J + \frac{1}{2}$ indicates that the height-varying eddy diffusivity $K(z)$ is to be evaluated on the interface separating the $(J+1)$ th and J th grid planes.

Specify grid-lengths $\Delta x \sim 0.5$ m, $\Delta z \sim 0.2$ m. For two choices $S_c = (1, 0.63)$ of the Schmidt number, compare your calculated solution $\bar{c}(100, z)$ at $x = 100$ m *on a graph* that also shows the observations (Table 1) of Project Prairie Grass run 57, for which the meteorological situation was $u_* = 0.50$ m s⁻¹, $z_0 = 0.0058$ m. (Note: to compare with Table 1, you’ll need to scale your computed concentrations the same way, that is, you multiply your solution for \bar{c} by $z_0 u_* / k_v$.)

¹The field of \bar{c} is the analog of the crosswind integrated concentration $\chi = \chi(x, z)$ due to a steady point source, and Project Prairie Grass provided field measurements of the latter.

Method

The above algorithm can be cast in the form

$$c_J \bar{c}_{I,J+1} + b_J \bar{c}_{I,J} + a_J \bar{c}_{I,J-1} = D_{I,J}, \quad J = 1..J_{\max} \quad (5)$$

where the c_J, b_J, a_J are the “neighbour coefficients” (following the naming scheme of our Matlab solver), and $\bar{c}_{I,J}$ is the concentration matrix. The term $D_{I,J}$ contains information (only) from the upstream column at $I - 1$, so we can regard it as a known. Thus we may frame our problem of finding the vertical column of values of \bar{c} at downstream location I (given the column at $I - 1$) as a matrix problem,

$$\mathbf{M} \mathbf{C} = \mathbf{D} \quad (6)$$

where \mathbf{M} is the coefficient matrix, and is tridiagonal. Thus we may find the unknown column (i.e. $\bar{c}_{I,J} \forall J$) as

$$\mathbf{C} = \mathbf{M}^{-1} \mathbf{D} \quad (7)$$

which in Matlab syntax is $\mathbf{C} = \mathbf{M} \setminus \mathbf{D}$.

In principle we need only once compute \mathbf{M}^{-1} , and once known we could use it repeatedly to step down the I axis: each time we get a new \mathbf{C} column matrix we recompute \mathbf{D} , and repeat the operation. Each step gives us \mathbf{C} at a column further down the x -axis by a distance Δx .

To compute the new \mathbf{C} column matrix for each step down the x -axis you may use either Matlab’s built-in matrix inversion $\mathbf{C} = \mathbf{M} \setminus \mathbf{D}$, or a Matlab implementation of the (less robust) TDMA method (“TDMA_solver.m”).

Upper and lower boundary conditions

Set the top of your domain $z(J_{\max})$ sufficiently high (say, at least 30 m) that $\bar{c}(J_{\max}) = 0$. Then your coefficients at $J = J_{\max}$ are

$$\begin{aligned} b_{J_{\max}} &= 1 \\ c_{J_{\max}} &= \text{not used} \\ a_{J_{\max}} &= 0 \\ D_{J_{\max}} &= 0 \end{aligned} \quad (8)$$

If we presume our gas does not react with the ground, we want zero flux to ground, which is assured by requiring $\bar{c}(I, 1) \equiv \bar{c}(I, 2)$. Thus at $J = 1$ the needed coefficients are

$$\begin{aligned} b_1 &= 1 \\ c_1 &= -1 \\ a_1 &= \text{not used} \\ D_1 &= 0 \end{aligned} \quad (9)$$

Inlet boundary condition

How is your solution going to “know” there is a source? Let J_h be the height index of the cell the physical source will lie within, and let the streamwise index value $I = 1$ correspond to a column of gridpoints aligned at the source location. The easiest approach is to set the inlet or inflow concentration profile as

$$\bar{c}(1, J) = \begin{cases} 0 & \text{if } J \neq J_h, \\ 1/(\bar{u}(J_h) \Delta z) & \text{if } J = J_h \end{cases} \quad (10)$$

which guarantees that the total mass flux across the first interior plane $I = 1$ will be

$$Q = \sum_J \bar{c}(0, J) \bar{u}(J) \Delta z = 1 \quad (11)$$

Table 1: Normalized (and dimensionless) crosswind-integrated concentration $z_0 u_* \chi / (k_v Q)$ observed at distance $x = 100$ m from the point source (height $h_s = 0.46$ m) in Project Prairie Grass run 57; units are χ , kg m^{-2} and Q , $[\text{kg s}^{-1}]$. (This is equivalent to the normalized concentration $z_0 u_* \bar{c} / (k_v q_\ell)$ at 100 m downwind from a crosswind line source at h_s , with \bar{c} , $[\text{kg m}^{-3}]$ and Q , $[\text{kg m}^{-1} \text{s}^{-1}]$.)

z [m]	$z_0 u_* \chi / (k_v Q)$
17.5	1.5E-6
13.5	6.6E-6
10.5	1.56E-5
7.5	3.51E-5
4.5	7.9E-5
2.5	1.25E-4
1.5	1.53E-4
1.0	1.62E-4
0.5	1.70E-4