# EAS 471 $3^{rd}$ Scored Assignment (20%) Due: 7 Apr. 2016

### Option B: Numerical simulation of an idealized windbreak flow

Suppose an infinitely long, porous windbreak of height H runs along the y-axis at x = 0, while the mean wind is oriented along the x-axis. We will suppose the surface layer is neutrally stratified, such that the profiles of mean wind speed and of eddy viscosity far upstream from the disturbance are

$$\overline{u}_0(z) = \frac{u_{*0}}{k_v} \ln \frac{z}{z_0} , \qquad (1)$$

$$K_0(z) = k_v \, u_{*0} \, z \tag{2}$$

 $(u_{*0} \text{ is the upwind friction velocity, } k_v = 0.4 \text{ is the von Karman constant, and } z_0 \text{ is the surface}$ roughness length; the subscript "0" denotes a property in the upwind, undisturbed region of the flow). We will further assume — however unrealistic this might appear<sup>1</sup>— that the disturbance to the flow does not alter the eddy diffusivity, i.e. at any point the local eddy viscosity  $K(x, z) \equiv K_0(z)$ . (For the purposes of the exercise, specific values for  $u_{*0}$ ,  $z_0$  and H are given below.)

The windbreak drags on the flow, and a plausible (and widely accepted) statement of momentum conservation reads:

$$\overline{u}\frac{\partial\overline{u}}{\partial x} + \overline{w}\frac{\partial\overline{u}}{\partial z} = \frac{-1}{\rho_0}\frac{\partial\overline{p}}{\partial x} - \frac{\partial u'^2}{\partial x} - \frac{\partial \overline{u'w'}}{\partial z} - k_r \overline{u} |\overline{u}| \,\delta(x-0) \,s(z-H) \,. \tag{3}$$

This is the steady-state, Reynolds-averaged  $\overline{u}$ -mtm equation, the final term being an empirical "momentum sink", localized at the windbreak:  $k_r$  is known as the "resistance coefficient" or "pressure loss coefficient", and is dimensionless. The localizing functions are the delta-function  $\delta(x-0)$  and a unit step function s(z-H) defined

$$s(z - H) = \begin{cases} 1, & z \le H \\ 0, & z > H \end{cases}$$
(4)

Notice that the product  $\overline{u} |\overline{u}|$  guarantees that the drag, proportional in magnitude to the square of the wind speed, always *opposes* the wind.

<sup>&</sup>lt;sup>1</sup>In fact it is not a bad first approximation. Wilson (1985) tested a range of "closure assumptions", including this one, i.e. an eddy viscosity closure using the undisturbed eddy viscosity  $K_0(z)$ . Simulations were compared against measurements, and this closure does capture the gross effect of the windbreak.

#### Numerical formulation

For our assignment we shall radically simplify Eqn. (3). Assuming the windbreak is very porous  $(k_r \ll 1)$  so that the disturbance to windspeed is small, we invoke a perturbation expansion

$$\overline{u}(x,z) = \overline{u}_0(z) + k_r \overline{u}_1(x,z), \qquad (5)$$

$$\overline{w}(x,z) = k_r \,\overline{u}_1(x,z) \,, \tag{6}$$

$$\overline{p}(x,z) = k_r \overline{p}_1(x,z) . \tag{7}$$

Then, retaining only terms to first order in  $k_r$ , one obtains

$$\overline{u}_0(z)\frac{\partial\overline{u}_1}{\partial x} = \frac{\partial}{\partial z}\left(K_0(z)\frac{\partial\overline{u}_1}{\partial z}\right) - \frac{1}{\rho_0}\frac{\partial\overline{p}_1}{\partial x} - \overline{u}\left|\overline{u}\right|\delta(x-0)s(z-H).$$
(8)

where the shear stress has been modelled using an (undisturbed) eddy viscosity  $K_0(z)$ .

The sum of the pressure gradient and the momentum sink can be represented by the following *effective* pressure field (Wilson et al. 1990),

$$\frac{\overline{p}_1}{\rho_0}(x,z) = \frac{-\overline{u}_{0H}^2}{2\pi} \left[ \tan^{-1}\left(\frac{H+z}{x}\right) + \tan^{-1}\left(\frac{H-z}{x}\right) \right]$$
(9)

in which  $\overline{u}_{0H}$  is the windspeed at H far upstream from the windbreak. Eq. (8), a linear PDE, combines elements of equations studied earlier in the course<sup>2</sup>. At the inflow boundary the wind speed perturbation is zero, but it is "driven" away from zero by the effective pressure disturbance; and it is subject to streamwise advection and vertical diffusion (with a heightvarying diffusivity). By taking a finite difference along the *x*-axis, Eq. (9) can be used to evaluate the pressure gradient wherever needed; however for any cell that spans the barrier one must set  $\partial \overline{p}_1 / \partial x = 0$ .

#### Discretization

Let I,J be the horizontal and vertical indices for a uniform grid, with the spacing between nodes being  $\Delta x, \Delta z$ . Define your z(J) as

$$z(J) = z_0 + (J-1)\Delta z$$
 (10)

<sup>&</sup>lt;sup>2</sup>In particular,  $\overline{u}_1$  plays a role similar to the mean concentration  $\overline{c}$  of the first scored assignment: it is advected by the same mean wind, and subject to vertical diffusion with an eddy diffusivity that differs from that we applied to concentration only by a constant factor  $S_c$ .

and your x(I) as

$$x(I) = X_0 + (I - 1)\Delta x$$
 (11)

so that J = 1 corresponds to  $z = z_0$  (where  $\overline{u}_1$  vanishes) and I = 1 corresponds to  $x = X_0$ , the inflow boundary where, again,  $\overline{u}_1$  vanishes. If  $J_{mx}$  indexes the highest row of gridpoints, then  $\overline{u}_1(I, J_{mx}) = 0$ . Stated verbally, with this setup the wind disturbance vanishes on the inflow boundary, and along the uppermost and lowermost gridplanes. As x is a one-way axis, a downwind boundary condition is not needed, i.e. this is a marching problem on the x-axis.

Application of the control volume method (integration of the equation throughout a control volume surrounding the node) gives the following discretization, valid at all *interior* gridpoints (1 < I and  $1 < J < J_{mx}$ ):

$$\Delta z \,\overline{u}_{0}(J) \,\left[\overline{u}_{1}(I,J) - \overline{u}_{1}(I-1,J)\right] = \Delta x \, K_{J+1/2} \frac{\overline{u}_{1}(I,J+1) - \overline{u}_{1}(I,J)}{\Delta z} - \Delta x \, K_{J-1/2} \frac{\overline{u}_{1}(I,J) - \overline{u}_{1}(I,J-1)}{\Delta z} - S_{I} \, \Delta z \left[\frac{\overline{p}_{1}}{\rho_{0}}(I+1/2,J) - \frac{\overline{p}_{1}}{\rho_{0}}(I-1/2,J)\right], \quad (12)$$

where the "switch"  $S_{\rm I}$  is defined

$$\mathcal{S}_{\mathrm{I}} = \begin{cases} 1 & \text{if } x(\mathrm{I}) \neq 0, \\ 0 & \text{if } x(\mathrm{I}) = 0 \text{ and } z_{\mathrm{J}} \leq H, \end{cases}$$
(13)

so as to "turn off" the (favourable) pressure drop that occurs across the barrier (were that retained, one would need also to retain the delta-function momentum sink). I have dropped the subscript '0' on the eddy viscosity, and the notation "J + 1/2" means that K is to be evaluated at  $z_{\rm J} + \Delta z/2$ . Recall that  $\bar{p}_1/\rho_0$  is given by Eq. (9).

We now cast this algorithm into the form

$$c_J \,\overline{u}_1(\mathbf{I}, \mathbf{J}+1) + b_J \,\overline{u}_1(\mathbf{I}, \mathbf{J}) + a_J \,\overline{u}_1(\mathbf{I}, \mathbf{J}-1) = D(\mathbf{I}, \mathbf{J}), \quad \mathbf{J} = 1..\mathbf{J}_{\mathrm{mx}}$$
 (14)

where the  $c_J, b_J, a_J$  are the "neighbour coefficients", and  $\overline{u}_1(I, J)$  is the windspeed perturbation. The coefficients are:

$$a_J = -\frac{\Delta x}{\Delta z} K_{J-1/2}, \qquad (15)$$

$$b_J = \Delta z \,\overline{u}_0(\mathbf{J}) + \frac{\Delta x}{\Delta z} \,K_{\mathbf{J}+1/2} + \frac{\Delta x}{\Delta z} \,K_{\mathbf{J}-1/2} \,, \tag{16}$$

$$c_J = -\frac{\Delta x}{\Delta z} K_{J+1/2}, \qquad (17)$$

and

$$D(\mathbf{I}, \mathbf{J}) = \Delta z \ \overline{u}_0(\mathbf{J}) \ \overline{u}_1(\mathbf{I} - 1, \mathbf{J}) - S_\mathbf{I} \ \frac{\Delta x \ \Delta z}{\rho_0} \ \frac{\overline{p}_1(\mathbf{I} + 1/2, \mathbf{J}) - \overline{p}_1(\mathbf{I} - 1/2, \mathbf{J})}{\Delta x} \ .$$
(18)

Solve Eqs. (9 and 12-18) numerically to obtain the solution field  $\overline{u}_1(x_{\rm I}, z_{\rm J})$  for a windbreak with height H = 2 m standing at x = 0 on a surface with roughness length  $z_0 = 0.01$ m. Let the upwind friction velocity  $u_{*0} = 0.25 \,\mathrm{m\,s^{-1}}$  and evaluate  $\overline{u}_{0\rm H}$  accordingly. Use a computational domain bounded by the ground plane ( $z = z_0$ ), an upper plane at  $z = z(J_{\rm mx}) =$ 20H, an inflow plane at  $x = X_0 = -10H$  and a downwind plane at  $x = x(I_{\rm mx}) = 30H$ . Choose your gridlengths as  $\Delta x = \Delta z = H/10$ .

On a single graph plot your computed transects of  $\overline{u}_1$  versus x/H along gridplanes close to these levels: z/H = (0.4, 0.8, 1.2, 2, 5). On the transect at  $z/H \approx 0.4$ , read off the *minimum* value  $\overline{u}_1^{\min}$  and, noting from Eq. (5) that

$$\frac{\overline{u}_1}{\overline{u}_0} \equiv \frac{1}{k_r} \frac{\overline{u}(x,z) - \overline{u}_0(z)}{\overline{u}_0(z)} , \qquad (19)$$

compare your  $\overline{u}_1^{\min}/\overline{u}_0$  with the theoretical value given by Wilson et al. (1990), viz.

$$\frac{1}{k_r} \left(\frac{\Delta \overline{u}}{\overline{u}_0}\right)^{\min} = \frac{-1}{(1+2k_r)^{0.8}} \approx -1.$$
(20)

## References

- Wilson, J.D. 1985. Numerical studies of flow through a windbreak. J. Wind Eng. Indust. Aero., 21, 119–154.
- Wilson, J.D., Swaters, G.E., & Ustina, F. 1990. A perturbation analysis of turbulent flow through a porous barrier. Q.J.R. Met. Soc., 116, 989–1004.