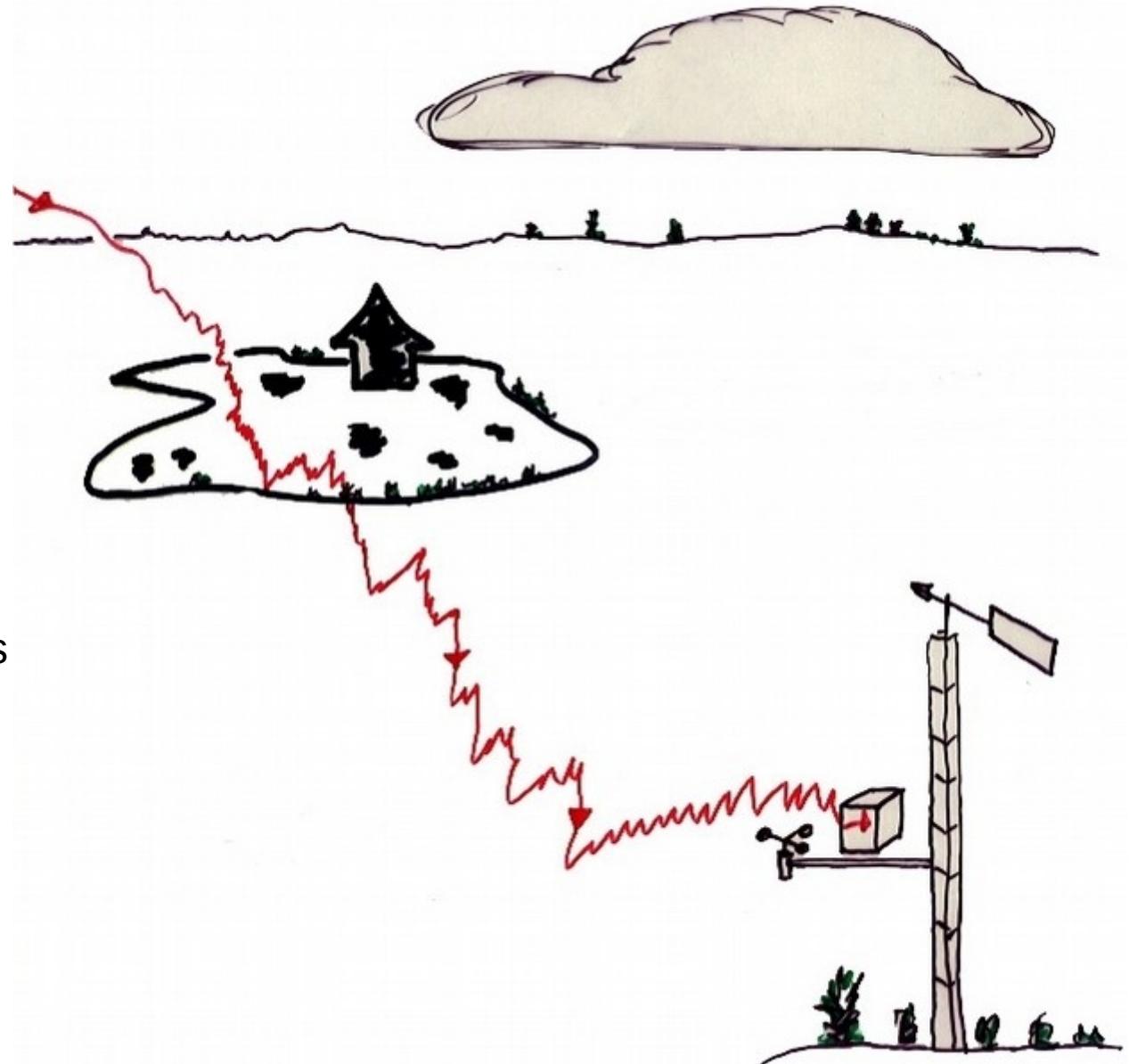
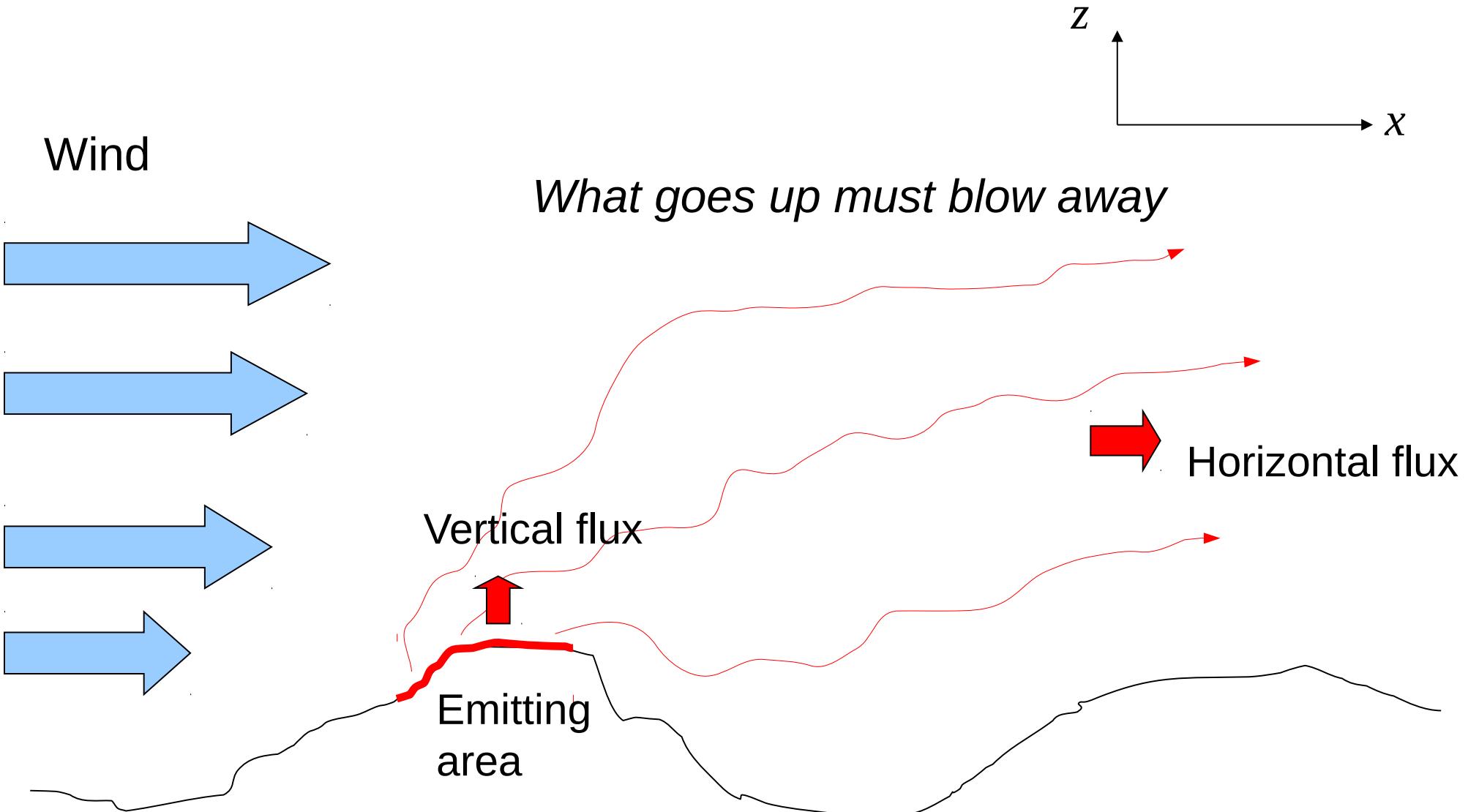


Micrometeorological methods to determine surface-air exchange

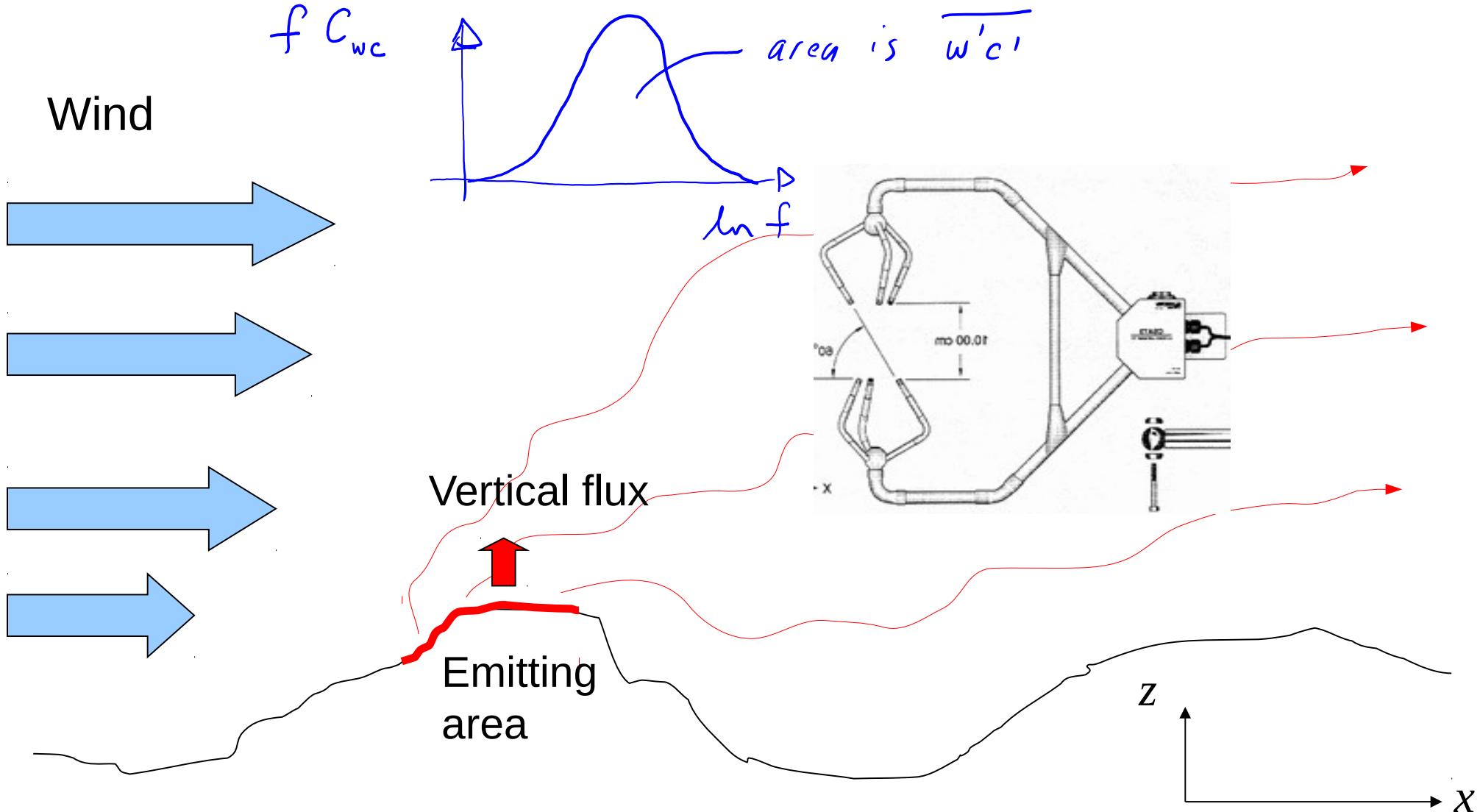
- Micromet methods – their basis
- The family of micromet methods
- Inverse dispersion in particular
- Applications

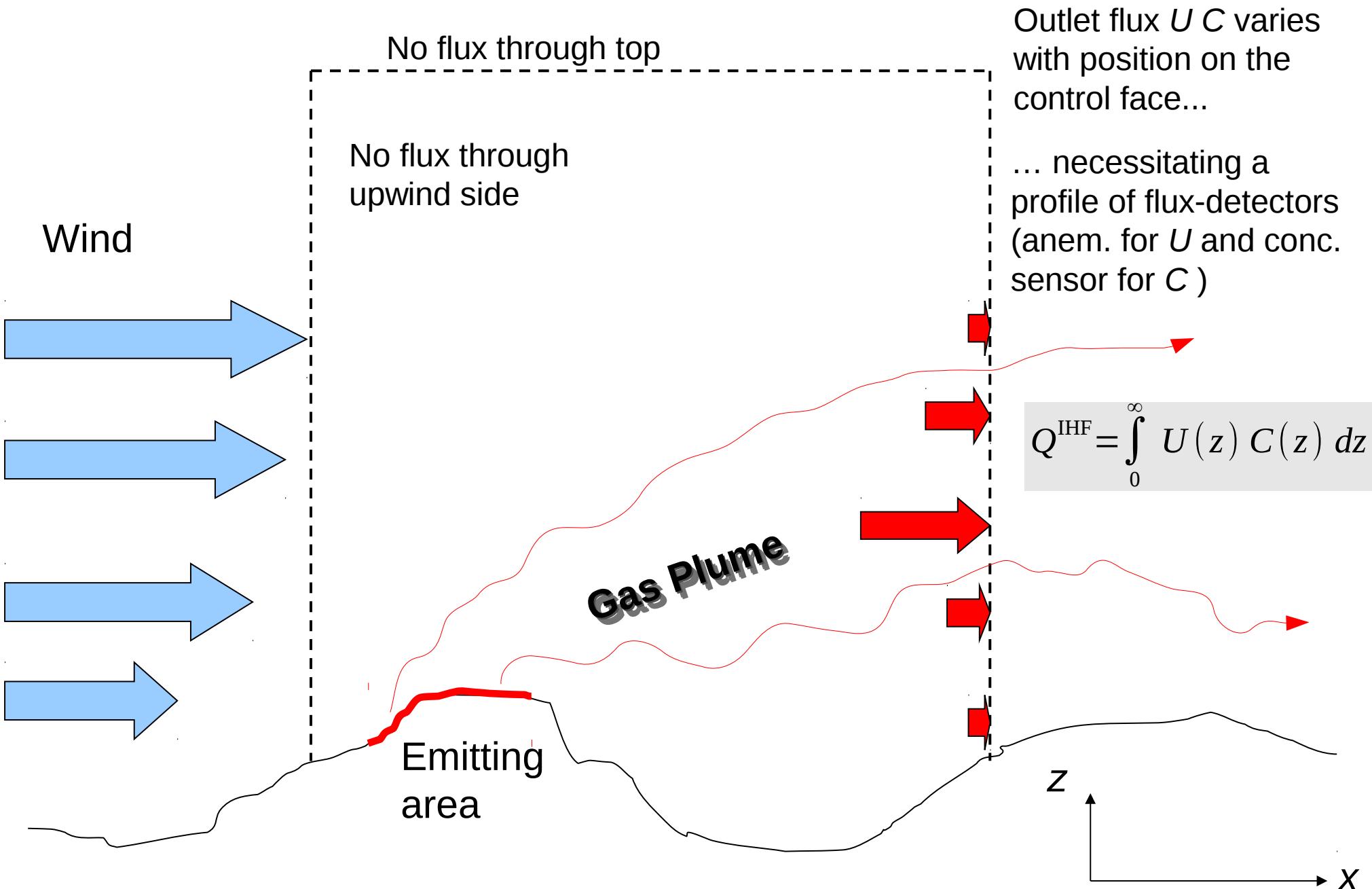


- as opposed to observing what goes missing from the substrate.
- winds and gas fluxes are averaged over periods of order 15 – 60 min.



- e.g. eddy covariance: fast gas detector measuring concentration $c(t)$ is paired with a sonic anemometer giving $w(t)$. Processing gives the vertical eddy flux $\overline{w'c'}$
- provisos on site suitability apply – requires large fetch of uniform source





- assume instruments are immersed in a constant flux layer,
i.e. $Q = \overline{w'c'}$ is independent of height between ground and z_2

- adopt gradient-diffusion model

$$Q = -K_c \frac{\partial C}{\partial z}, \quad \overline{w'\theta'} = -K_h \frac{\partial \bar{\theta}}{\partial z}$$

- assume the (kinematic) heat flux $\overline{w'\theta'}$ is known (e.g. measured by a sonic anemometer)
- then

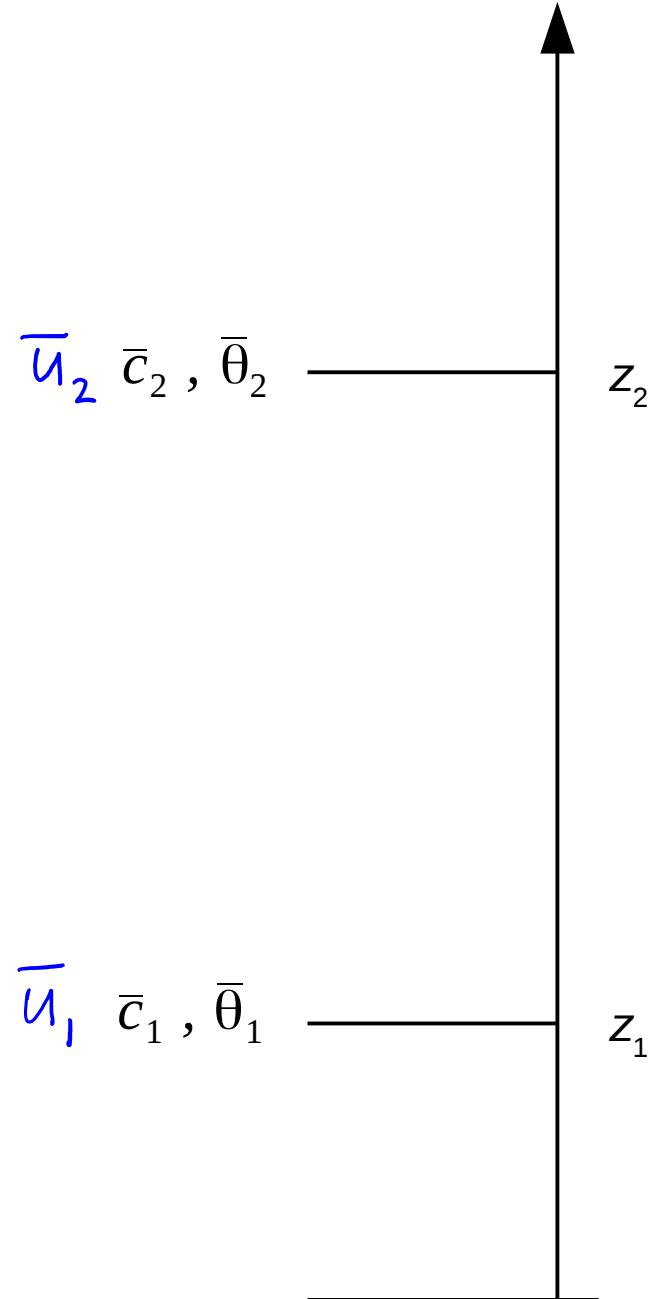
$$\frac{Q}{\overline{w'\theta'}} = \frac{K_c}{K_h} \frac{\bar{c}_2 - \bar{c}_1}{\bar{\theta}_2 - \bar{\theta}_1}$$

- or, assuming MOST applies

$$Q = \overline{w'\theta'}, \quad \frac{\phi_h(z_m/L)}{\phi_m(z_m/L)} \frac{\bar{c}_2 - \bar{c}_1}{\bar{\theta}_2 - \bar{\theta}_1}$$

where $\overset{\text{mean}}{z_m}$ is often specified as $(z_1 z_2)^{1/2}$.

- as an alternative to measuring the heat flux and temperature profile, may measure the wind profile and take u_*^2 as the companion flux



$$\overline{u'w'} = -u_*^2 = -K_m \frac{\partial \bar{u}}{\partial z}$$

$$\frac{Q}{-u_*^2} = - \frac{K_c}{K_m} \frac{\bar{c}_2 - \bar{c}_1}{\bar{u}_2 - \bar{u}_1} = - \frac{\phi_m(z_m/L)}{\phi_c(z_m/L)} \frac{\Delta \bar{c}}{\Delta \bar{u}}$$

mtm

Best fit MO profiles of \bar{u} and $\bar{\theta}$ to our

$$\bar{u}_1, \bar{u}_2, \bar{\theta}_1, \bar{\theta}_2$$

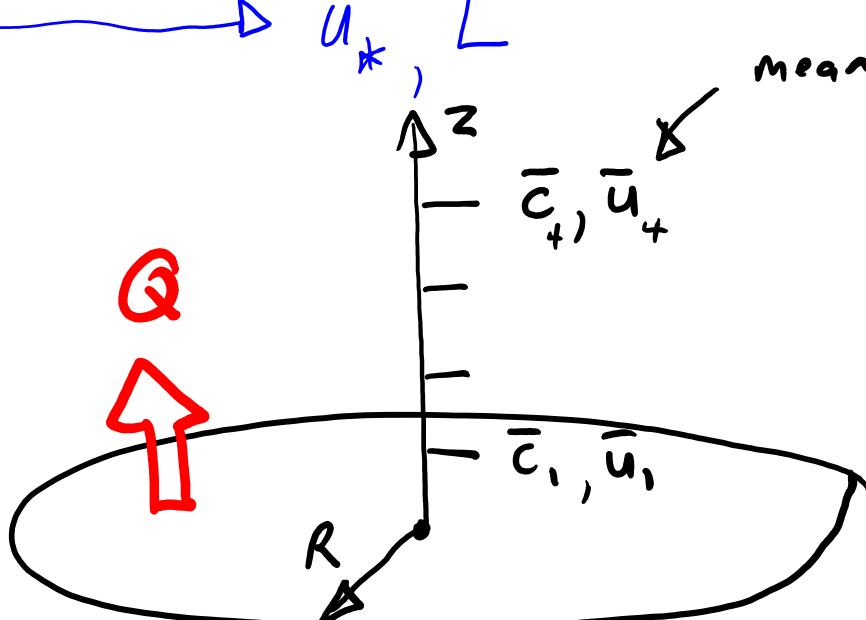


$$u_*, L$$

mean cup windspeed

$$QR = \int_0^\infty \bar{u}(z) \bar{c}(z) dz$$

(a variant of IHF)

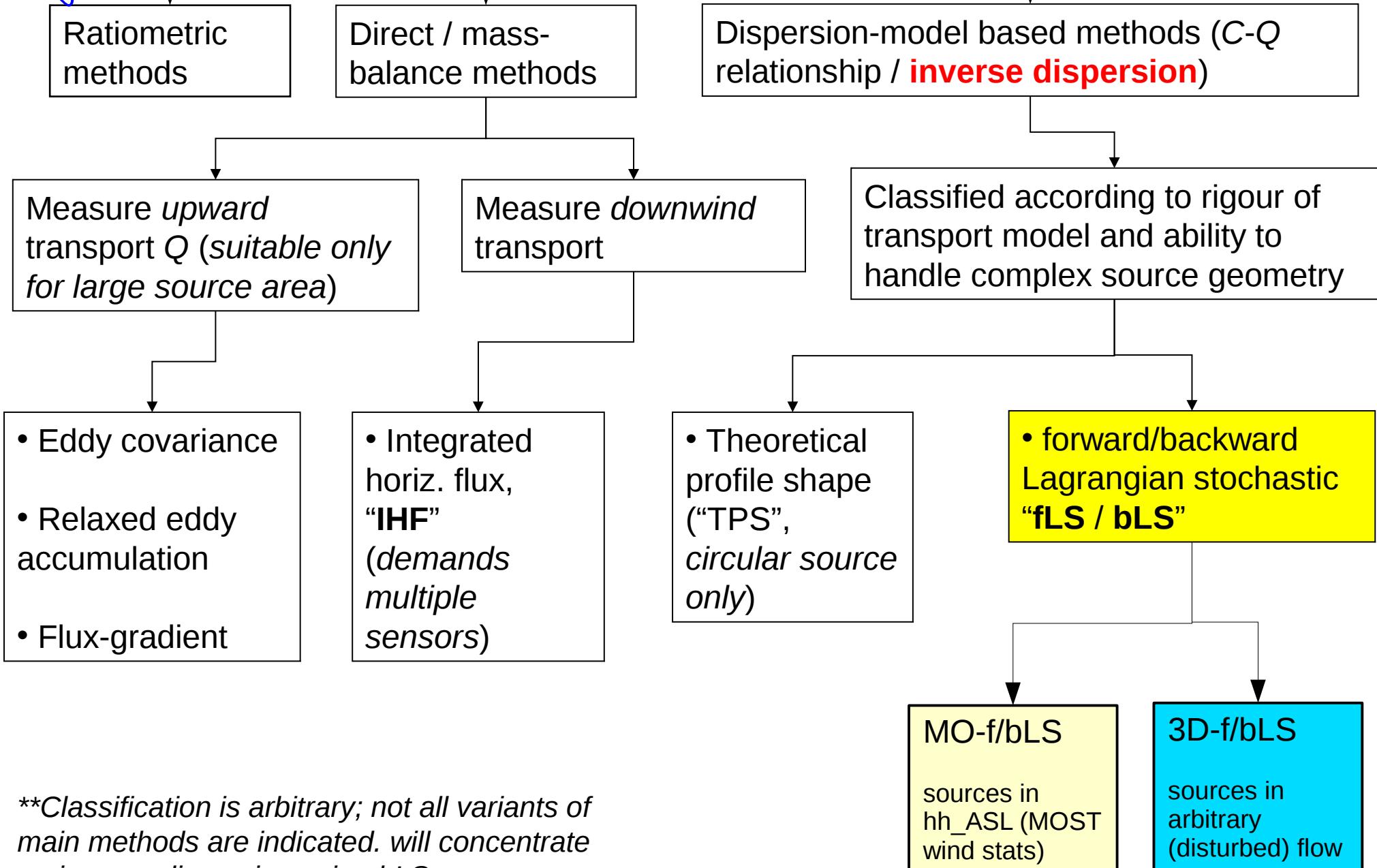


TPS method reduces

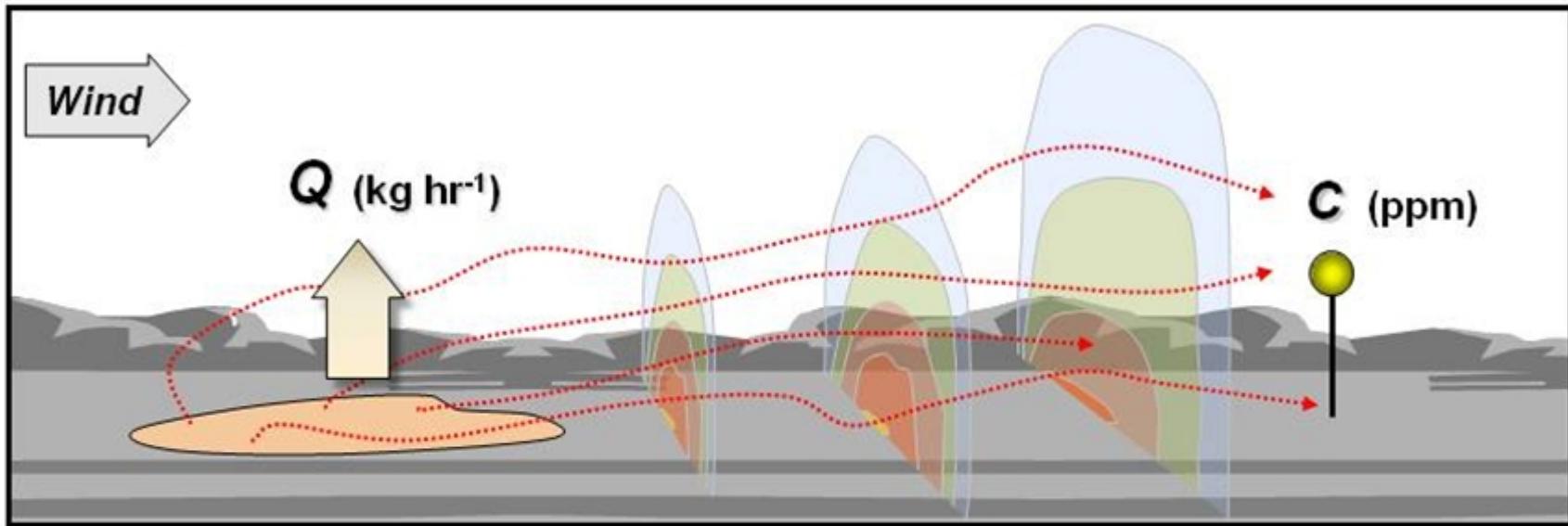
expt' input to a single

\bar{u} and single \bar{c} at a special height $Z^{INST.}$

entails using a tracer gas with known
release rate



**Classification is arbitrary; not all variants of main methods are indicated. Will concentrate on inverse dispersion using bLS

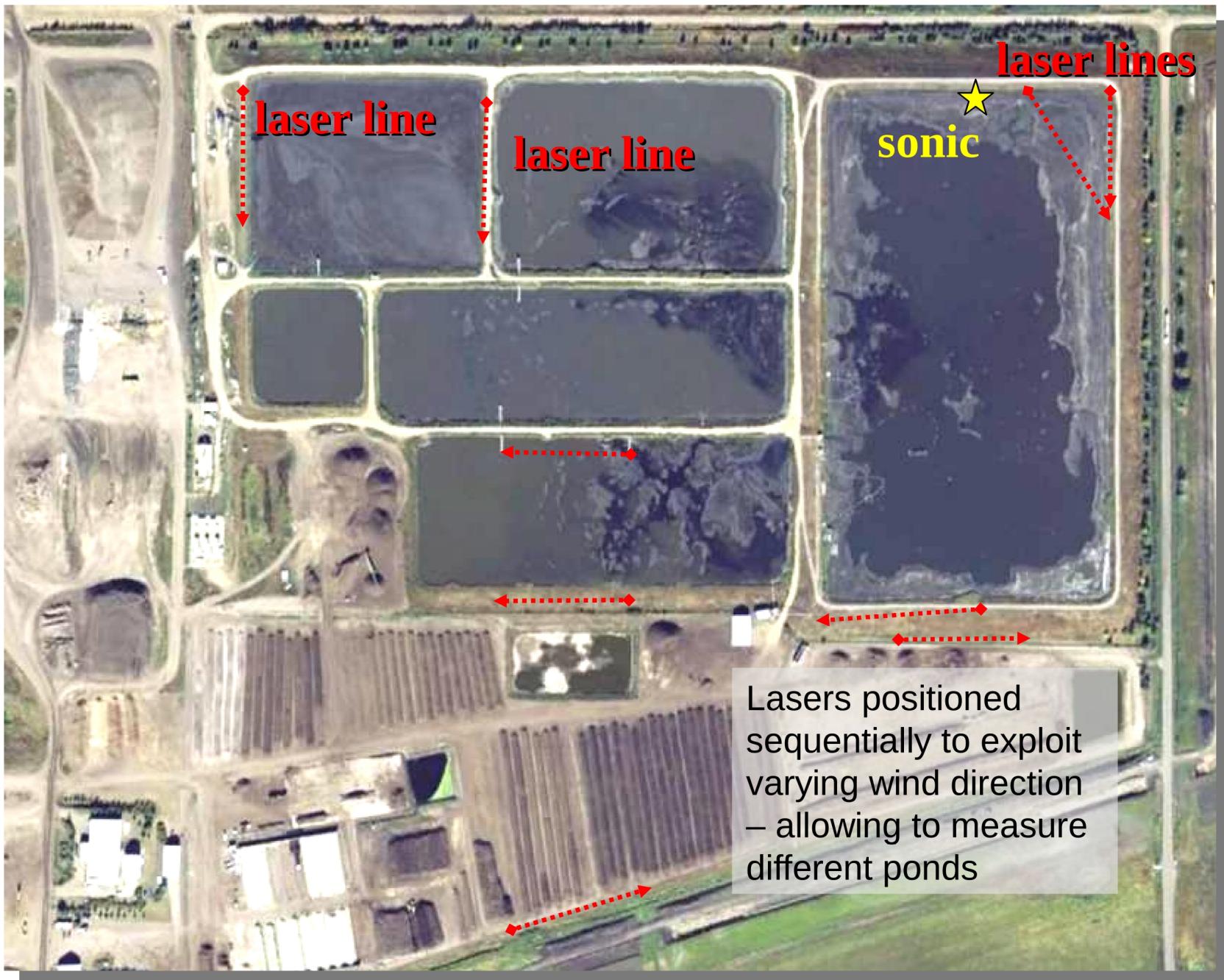


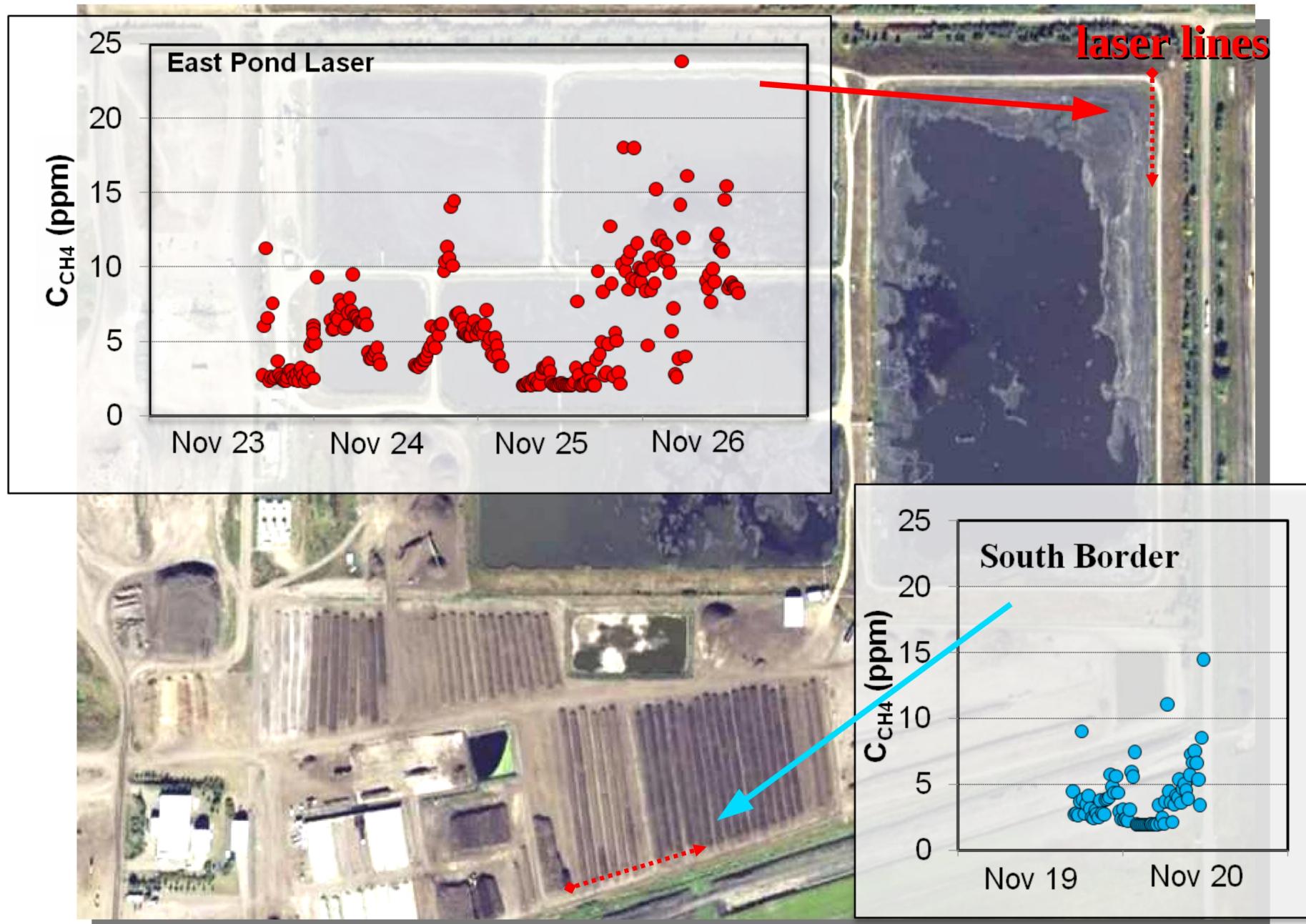
- Atmospheric dispersion model relates downwind concentration C to emission rate Q for prevailing regime of wind & turbulence
- Measurement of C (minus background) + model permits to infer Q
- Approach blends data + theory
- “WindTraX” is a Lagrangian stochastic (LS) particle trajectory model appropriate for inverse dispersion on the surface layer scale – assumes wind statistics obey MOST





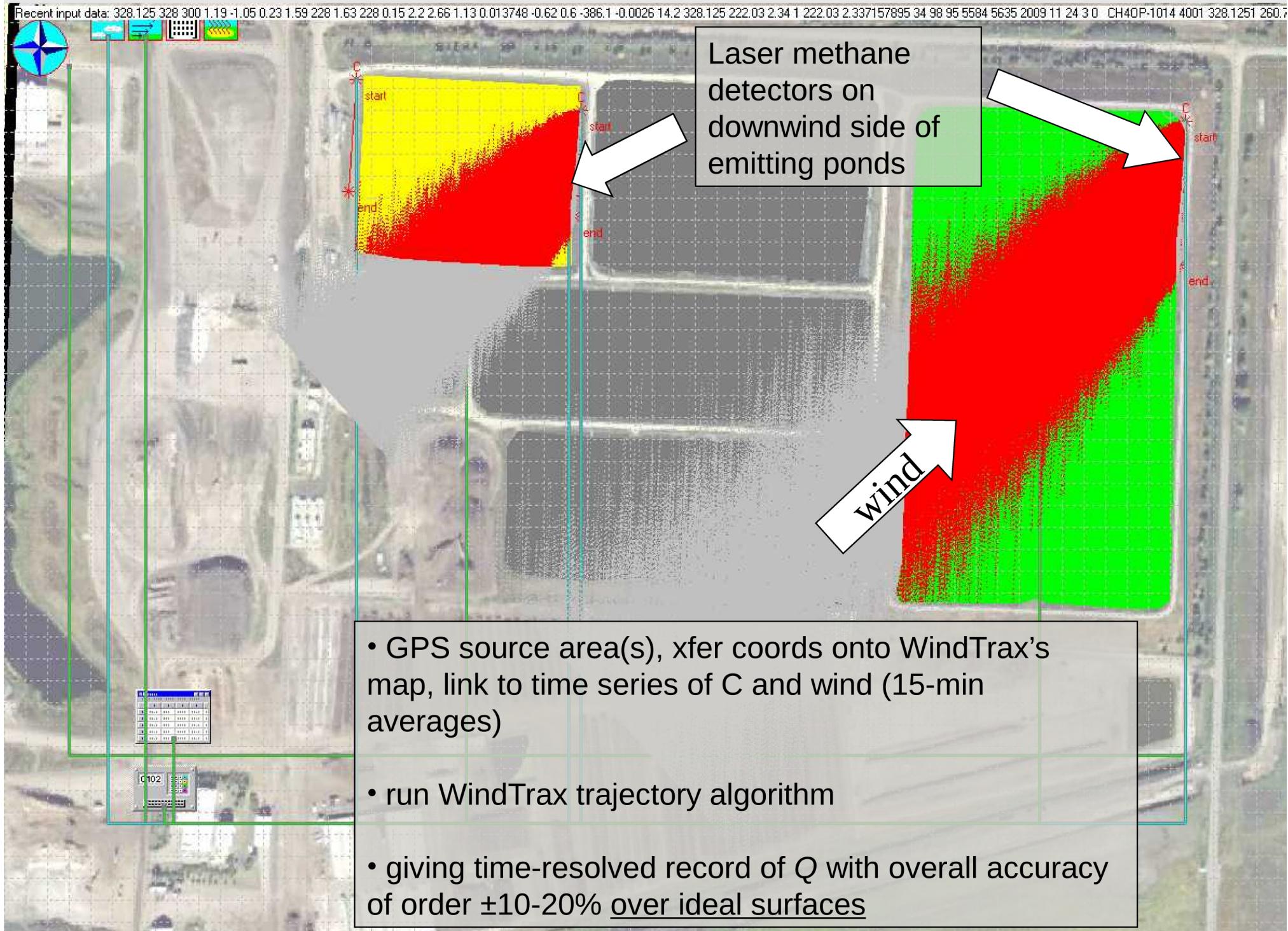
Line-averaging laser methane detector
(pathlength typically ~ 100 m)

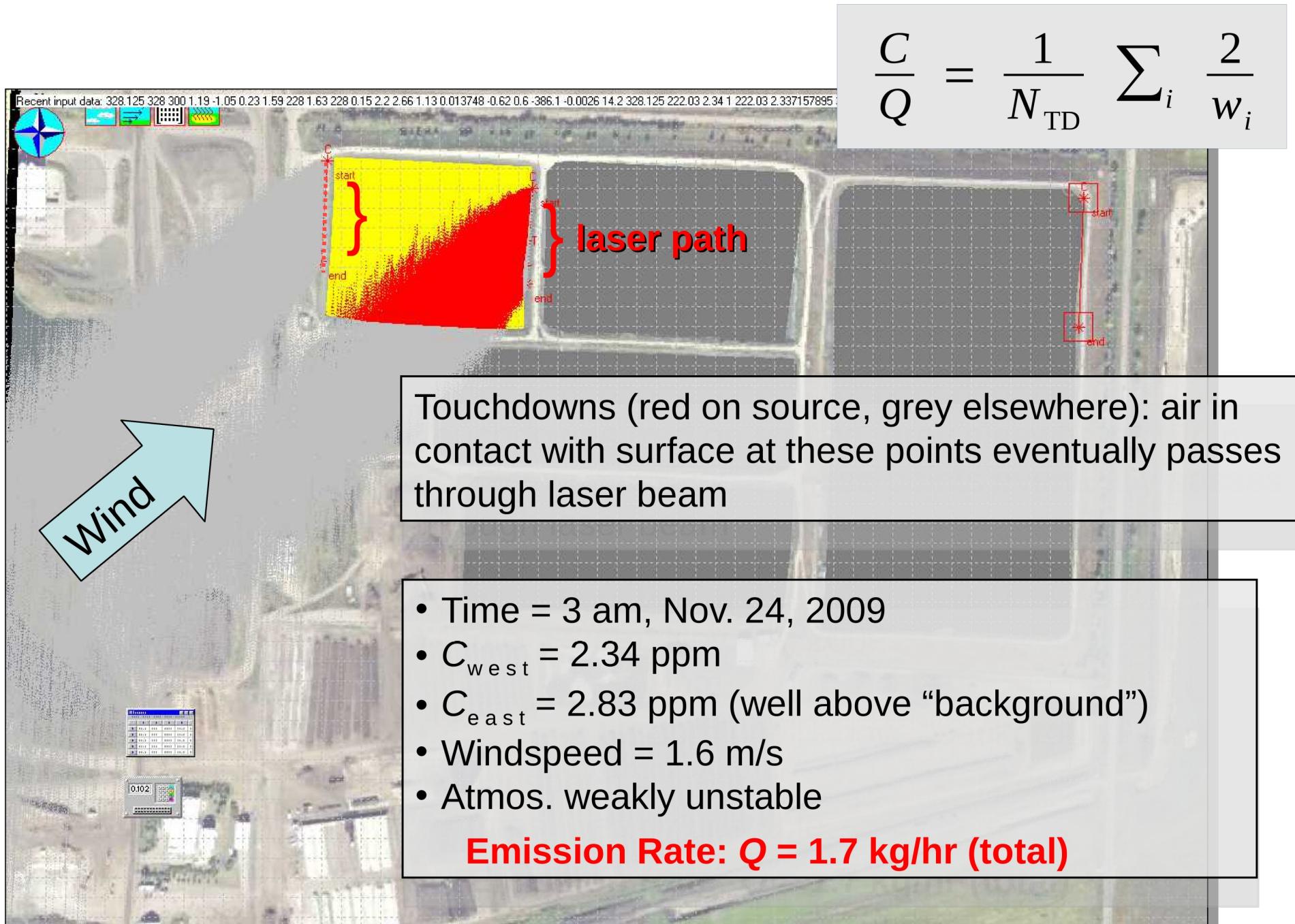


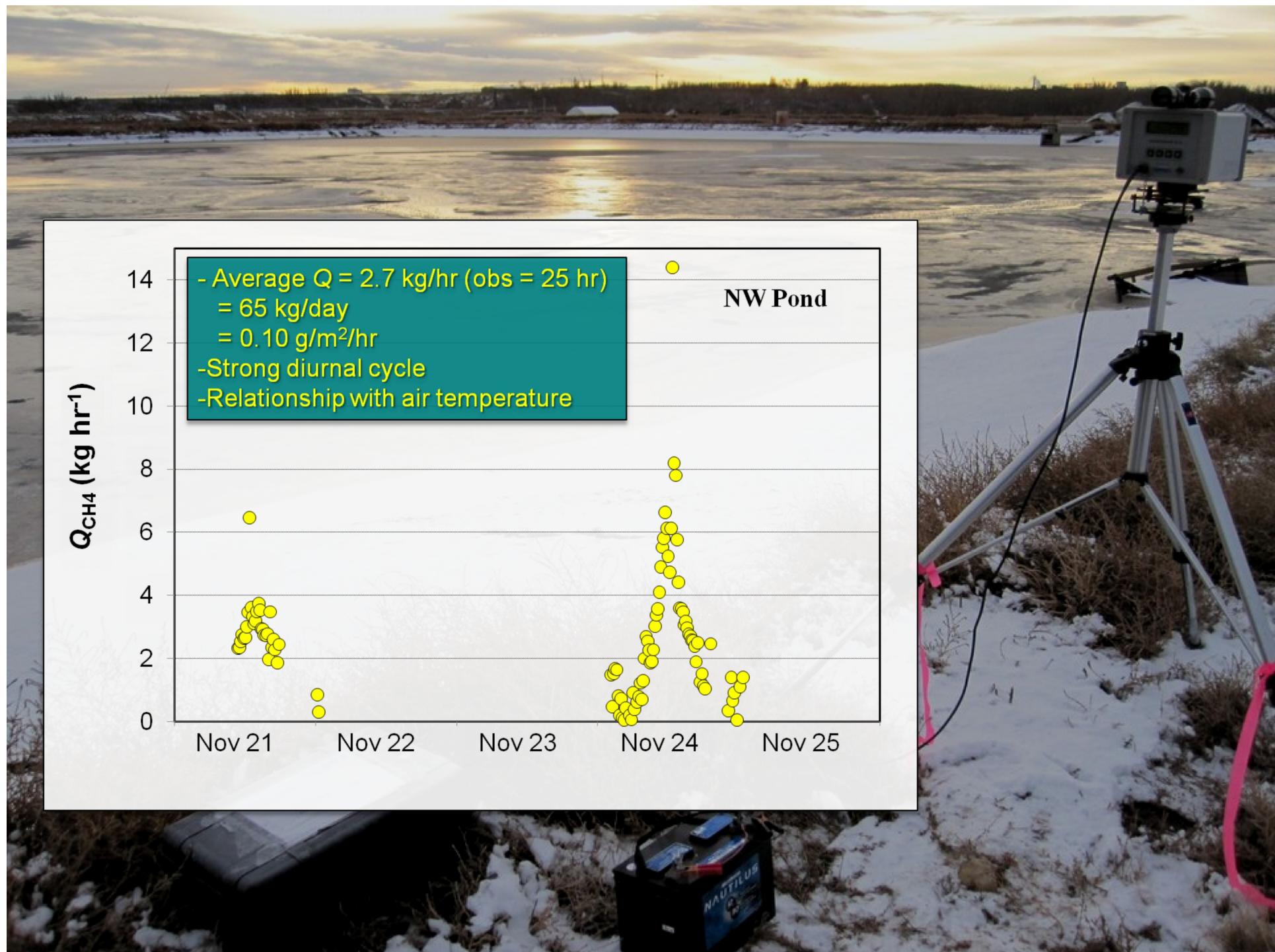


WindTrax facilitates MO-bLS – here showing “touchdown clouds”

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CH4 emission rate by eddy covariance & inverse dispersion (Casandra Brown) 14b









Feedlot, CH4 & NH3 emission rate vs DOY

