

Exercise – the steady state surface layer TKE budget according to MOST

$$\frac{\partial k}{\partial t} = 0 = \underbrace{-\overline{u'w'} \frac{\partial U}{\partial z} - \overline{v'w'} \frac{\partial V}{\partial z}}_{\text{Shear production, } \mathbf{P}_s} + \underbrace{\frac{g}{\theta_0} \overline{w'\theta'}}_{\text{Buoyant production, } \mathbf{P}_B} - \underbrace{\epsilon}_{\text{Dissipation (conversion of TKE to heat)}} - \underbrace{\frac{\partial}{\partial z} \overline{w' \left(\frac{p'}{\rho} + \frac{u'u' + v'v' + w'w'}{2} \right)}}_{\text{Turbulent \& pressure transport, } \mathbf{T}_T}$$

Assume the coordinate system is chosen to make $V=0$ throughout the surface layer, and neglect the term T_T (known as the assumption that the TKE budget is in "local equilibrium"). Adopting the assumptions and/or definitions inherent to MOST (including the definition of the Obukhov length L), prove that the TKE dissipation rate can be written

$$\frac{k_v z \epsilon}{u_*^3} = \phi_\epsilon \left(\frac{z}{L} \right) \quad \text{where}^{**} \quad \phi_\epsilon \left(\frac{z}{L} \right) = \phi_m \left(\frac{z}{L} \right) - \left(\frac{z}{L} \right)$$

$$^{**} \quad \phi_m \left(\frac{z}{L} \right) = \frac{k_v z}{u_*} \frac{\partial U}{\partial z} \quad \text{is the MO function for the dimensionless wind shear}$$